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1 *Mineral phases in metakaolin-portlandite pastes cured 15 years at 60°C. New data for*
2 *scientific advancement*

3

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13 **Abstract**

14 One of the problems addressed by the scientific community in connection with cement
15 matrix hydration is the evolution of hydrated phases under certain, primarily
16 temperature-related, curing conditions. Such conditions are of utmost importance when
17 the cement generates metastable hydrated phases that develop into stable phases,
18 inducing substantial physical-mechanical changes in the matrix. One such instance
19 arises during the pozzolanic reaction between metakaolinite and calcium hydroxide at
20 60 °C when metastable hexagonal phases co-exist with cubic stable phases. Such
21 processes are well understood in the short and medium term (<5 years). Evolution at
22 longer curing times has not been studied, however. This paper describes the evolution of
23 the initial (stable and metastable) mineral phases in the metakaolinite / Ca(OH)₂
24 reaction toward a zeolite-like phase after 15 year curing at 60 °C. An understanding of
25 such mineralogical changes is scientifically significant, given their direct impact on the
26 engineering properties of cement-based matrices.

27 **Keywords:** Metakaolinite, pozzolanic reaction, metastable phases, temperature,
28 mineralogical evolution, zeolites.

30 1. INTRODUCTION

31 For decades, the inclusion of calcined aluminosilica pozzolans in cement matrices
32 (Gbozee et al., 2018) has prompted a fair degree of uncertainty in connection with the
33 possible formation of metastable phases. Under some curing conditions such phases
34 may evolve into stable cubic phases, generating greater volume and porosity, with the
35 concomitant adverse effects on blended cement mechanical performance and durability
36 (Silva and Glasser, 1993).

37 Research has focused on the use of kaolinite (K)-based clay minerals which when
38 heated to temperatures of 400 °C to 700 °C originate metakaolinite (MK) (Kakali et al.,
39 2001; Fabbri et al., 2013; Bich et al., 2009). Further to research conducted by Ptacek et
40 al. (2014), K converts to MK in three stages: lamination, dehydroxylation and MK
41 formation. The sequential order of the first two depends directly on the activation rate.
42 As MK exhibits high pozzolanicity, it is ideal for use as a supplementary cementitious
43 material in commercial blended cement manufacture (Siddique and Klaus, 2009;
44 Rashad, 2013) as acknowledged in international standards (European Standard, 2011).

45 A number of authors (Murat, 1983; Serry et al., 1984; Santos et al., 2014; Frías and
46 Cabrera, 2001) have shown that the main hydrated phases resulting from the pozzolanic
47 reaction in pure MK/lime systems at ambient temperature are C-S-H gels, C_4AH_{13} ,
48 C_2ASH_8 (strätlingite) and $C_4A_3CH_{11}$. The amount of each phase formed depends on
49 factors such as the MK/lime and MK/binder ratios, hydration time and curing humidity.
50 The implications of the presence of these hydrated phases in the cementitious matrix
51 include a decline in microporosity (Frías and Cabrera, 2000; Mleza and Hajjaji, 2012;
52 Sabir et al., 2001; Jiang et al., 2015), which has a direct impact on physical-mechanical
53 properties and durability. Higher curing temperatures, in turn, are known to favour the

54 formation of the (variable composition) cubic hydrogarnet phase ($C_3AH_6-C_3AS_3H_n$)
55 (Klimesch and Ray, 19989; Rios et al., 2009; Cabrera and Frías, 2001; Martínez and
56 Frías, 2011).

57 The study of the effect of temperature on pozzolanic reactions in cement matrices dates
58 back to the nineteen nineties, when Silva et al. (1993) reported that in high-lime
59 matrices C-S-H gels and hydrogarnet are stable up to temperatures of 55 °C. In contrast,
60 metastable phases C_2ASH_8 and C_4AH_{13} destabilise at that temperature, ultimately
61 forming a stable cubic hydrogarnet. Those developments induce a decline in matrix
62 volume of nearly 13 %, with a rise in porosity and concomitant loss of strength.
63 Amboise and Murat (1985) established the possible chemical reactions for phase
64 formation and evolution, as follows:

65



68

69 With a view to obtaining a fuller understanding of these reactions at longer times in
70 pure MK/lime systems, Martínez and Frías (2011), Frías and Cabrera (2002), Frías and
71 Sánchez de Rojas (2003) and Frías et al. (2003) studied a (1:1 by weight) MK/lime
72 system at 60 °C for 123 d. According to their findings, metastable hexagonal phases co-
73 existed with a stable cubic phase resulting from the pozzolanic reaction in a medium
74 with traces of $Ca(OH)_2$. At that curing age (123 d), nearly all the metastable phases had
75 evolved into stable hydrogarnet phases, with a perceptible decline in C-S-H gel
76 concentration. Pesce et al. (2014) reported similar findings for 28 day MK/lime blends
77 cured at 25 °C, with the formation of C-S-H gels, C_4AH_{13} and C_2ASH_6 . These authors

78 also observed a higher concentration of cubic phase C_3ASH_6 and a lower C-S-H gel
79 content relative to other systems studied.

80 At longer times (5 years), Frías (2006a) showed hexagonal metastable C_2ASH_8 and
81 stable cubic hydrogarnet to be the predominant crystalline hydrated phases in a
82 $Ca(OH)_2$ -free MK/lime system (1:1 by weight), along with traces of vertumnite
83 ($Ca_4Al_4Si_4O_6(OH)_{24} \cdot 3H_2O$), a strätlingite polytype differing from the general form in
84 its silicon and water content (Rinaldi et al., 1990).

85 Whilst today's civil works are designed to a service life of over 100 years, no scientific
86 studies on the evolution of these MK/lime mineralogical phases under such extreme
87 curing conditions at longer curing times have yet been forthcoming. This article
88 describes a first-time study of the mineral phases present in MK/lime systems cured at
89 $60^\circ C$ for 15 years. The findings, which further the scientific understanding of the
90 pozzolanic reaction in aluminosilicate systems, will have a direct impact on the
91 properties of cementitious systems and consequently their possible applications in
92 engineering.

93

94 **2. EXPERIMENTAL PROCEDURES**

95 **2.1 Materials**

96 A commercial metakaolin (MK) manufactured in the UK was used in this study. The
97 XRF determined elemental composition in Table 1 shows that the MK used consisted
98 primarily in Si and Al. Its BET specific surface was found to be $15\text{ m}^2/\text{g}$. XRD analysis
99 (spectrum not shown) revealed the presence of traces of quartz and mica. Analytical
100 grade lime $Ca(OH)_2$ was used.

101

102 **2.2 Blended MK- $Ca(OH)_2$ pastes**

103 The blended pastes were prepared by mixing a 1:1 (by weight) ratio of MK and lime
104 with water at a water/solid ratio of 2.37. The resulting slurry was placed in a lidded
105 plastic airtight container and submerged in water at 60 °C for 15 years as depicted in the
106 sketch in Figure 1.

107 After the 15 year curing, the solid paste was removed from the container, mixed with
108 acetone and heated in a laboratory oven at 40 °C for 24 h to detain the pozzolanic
109 reaction.

110

111 **2.3 Instrumental techniques**

112 The mineralogical composition of the bulk samples was determined by X-ray powder
113 diffraction (XRD) on a Siemens D-5000 (Munich, Germany) X-ray diffractometer fitted
114 with a Cu anode, as well as a 2.0 mm divergence and a 0.6 mm reception slit. The
115 operating settings were 30 mA, 40 kV; scanning at 0.041° (2θ) steps; count time, 3 s.
116 Rietveld analysis was performed using rutile as the internal standard to quantify the
117 amorphous phase (Rietveld, 1969; Ruan and Ward, 2002). The quantification findings
118 also specified the X^2 goodness of fit, which compared peak intensities calculated from
119 the spectrum to the experimental values based on phase weight. Although ideally X^2
120 should be equal to 1, since the discrepancies between the observed and calculated data
121 widen substantially in materials containing over three phases, values of $X^2 < 10$ are
122 generally accepted as indicative of valid results.

123 Sample morphology and microanalysis were determined on a Thermo Fisher Scientific
124 (FEI Company) Inspect scanning electron microscope (Hillsboro, OR) equipped with an
125 X-ray dispersive analyser. Chemical composition was computed as the average of ten
126 analyses per sample, ± the resulting standard deviation. The findings are expressed in
127 wt% of oxides.

128 Raman spectra were recorded on a Renishaw inVia confocal Raman spectroscope fitted
129 with a Leica microscope and an electrically refrigerated CCD camera. The instrument's
130 diode laser delivered a 785 nm excitation frequency and a 25 mW laser beam that was
131 focused on a quartz tube containing the sample. Silicon was used to calibrate the
132 frequencies.

133 The ^{27}Al and ^{29}Si MAS NMR spectra were recorded with a Bruker MLS-400
134 spectrophotometer operating at 79.49 MHz. The recording conditions for the ^{27}Al
135 spectra were: spinning speed, 10 kHz; pulse width, 2 μs ; relaxation delay, 5 s; No.
136 scans, typically 360. The conditions for the ^{29}Si spectra were spinning speed, 10 kHz;
137 pulse width, 5 μs ; relaxation delay, 20 s; No. scans, typically 1120. Tetramethylsilane
138 was used as a reference for the ^{29}Si and a 0.1 M aqueous solution of $\text{Al}(\text{NO}_3)_3$ for the
139 ^{27}Al chemical shifts. The 600 ^{27}Al CP-MAS scans made to generate the spectrum were
140 recorded at a contact time of 10 ms.

141 Fourier transform infrared (FTIR) analysis was conducted on a Thermoscientific Nicolet
142 6700 (Waltham, USA) spectrometer. One-mg samples pressed into 300 mg of KBr were
143 analysed in spectral range 4000 cm^{-1} to 400 cm^{-1} .

144 Thermogravimetric and differential thermal analyses (TG/DTA) were run on a Stanton
145 STA 781 instrument. Samples weighing 30 mg to 35 mg were heated at a rate of 10
146 $^{\circ}\text{C}/\text{min}$ in a nitrogen atmosphere.

147 **3. RESULTS**

148 **3.1 XRD-determined mineral phases**

149 The 15 year XRD spectrum for $\text{MK}/\text{Ca}(\text{OH})_2$ exhibited diffraction lines for three
150 CaCO_3 polymorphs (calcite, aragonite and vaterite) and gismondine (Figure 2). Calcite
151 was identified in the calcium carbonate group by reflections at 3.85 \AA ($23.1\ 2\theta$); 3.03 \AA

152 (29.4 2 θ); 2.49 Å (35.9 2 θ); 2.28 Å (39.5 2 θ) 2.09 Å (43.2 2 θ); 1.92 Å (47.3 2 θ) and
153 1.87 Å (48.6 2 θ); aragonite at 3.39 Å (26.3 2 θ); 3.27 Å (27.2 2 θ); 2.73 Å (32.8 2 θ);
154 2.70 Å (33.12 θ); 1.81 Å (50.4 2 θ); 1.74 Å (52.5 2 θ) and 1.72 Å (53.2 2 θ); and vaterite
155 primarily at 4.20 Å (21.1 2 θ) and 3.57 Å (24.9 2 θ). Gismondine was identified by
156 reflections at 9.98Å (8.85 2 θ); 7.26 Å (12.14 2 θ); 5.95 Å (14.85 2 θ); 5.74Å (15.4 2 θ);
157 5.01 Å (17.69 2 θ); 4.90 Å (18.1 2 θ); 4.45 Å (19.9 2 θ); 4.26 Å (20.82 2 θ); 3.36 Å (26.5
158 2 θ); 3.18 Å (28.01 2 θ); 3.12 Å (28.58 2 θ); 2.74 Å (32.62 2 θ); 2.66 Å (33.64 2 θ) and
159 206 Å (43.9 2 θ).

160

161 The Rietveld quantification of the mineral phases yielded the following results:
162 gismondine 30 %, aragonite 20 %, calcite 20 %, vaterite 6 % and amorphous phase
163 24 % ($X^2 = 3.7$). The sharp, narrow reflections for gismondine, aragonite and calcite
164 denoted a high degree of crystallisation. The presence of possible traces of non-
165 crystalline strätlingite and hydrogarnet could also be inferred from the low humps in the
166 baseline in the areas of the spectrum associated with their characteristic reflections:
167 respectively 12.6 Å (7.1 2 θ)) and 2.8 Å (31.9 2 θ). These two phases have been
168 identified as the main pozzolanic reaction products in 5 year materials (Frías, 2006a).

169

170 **3.2 SEM/EDX-determined mineral phases**

171 The elemental analysis of the various morphologies revealed by the textural study is
172 given in Table 2. Further to the combined SEM/EDX information, the morphologies
173 and mineral phases present were as follows.

174 Hexagonal plates and some crystallisation were observed on the conglomerate, which
175 served as a substrate (Table 2 and Figure 3). The plate surfaces were eroded and highly
176 porous while their edges were poorly defined and rounded, possibly denoting incipient
177 degradation. EDX microanalysis (Table 2) identified the plates as metastable phases
178 C_4AH_{13} (Figure 3a) and strätlingite (Figure 3b).

179 The conglomerates observed in the sample exhibited scanty defined morphologies:
180 skeletal forms with rounded edges and pock-holed surfaces (Figures 3 to 5 and Table 2).
181 They consisted primarily in silica and alumina with smaller proportions of sodium,
182 potassium and magnesium, with substantial quantities of randomly scattered calcium.
183 The appearance and composition of this sand were indicative of time-induced
184 degradation of the other materials present, such as C-S-H gels, C_4AH_{13} , strätlingite and
185 hydrogarnets.

186 Substantial numbers of fibres with a composition similar to that of C-S-H gel were
187 observed to form on the conglomerates (Figure 4a and Table 2). Calcium feldspar
188 crystals also formed (Figure 4b and Table 2), with smooth surfaces and well-defined
189 grain edges as a result of heterogeneous nucleation on the conglomerate substrate.

190 The conglomerate also served as a crystallisation substrate for calcium carbonates and
191 cubic calcium zeolites of the gismondine/garronite family. The calcium carbonate
192 crystals were the result of one of two processes: a) neo-formation of calcium carbonates
193 characterised by smooth surfaces and sharp grain edges (Figure 4c) and b) deterioration
194 of the calcium carbonates present in the starting paste, attested to by the rough surfaces
195 and poorly defined edges of the crystals (Figure 4d).

196 Gismondine, a calcium zeolite, exhibited a pseudo-cubic morphology ($P2_1/c$ symmetry),
197 pseudo-octahedral clustering and a Si/(Si+Al) ratio of close to 0.5 (Fischer and
198 Schramm, 1971). Its grain edges were well defined and its surfaces smooth,

199 indicating neo-formation (Figure 5a and Table 2). The calcium garronite in turn
200 (Figure 5b) consisted in compact radial conglomerates with Si/(Si+Al) ratios of over
201 0.5, rough surfaces and poorly defined crystal edges. Such poorly defined edges did not
202 denote degradation in this case, however, for garronite initially generates a disorderly
203 phase that over time tends to partial order and a scanty defined morphology. Its
204 structure exhibited four orientation domains and $I4_1/amd$ symmetry (Gottardi et al.,
205 1974) (Figure 5b and Table 2).

206 **3.3 Micro-Raman-determined mineral phases**

207 The Raman spectrum for the MK/lime sample (Figure 6) contained signals
208 characteristic of calcite (1087 cm^{-1} , 712 cm^{-1} and 280 cm^{-1}) and three wide signals at
209 990 cm^{-1} , 660 cm^{-1} and 445 cm^{-1} , attributable to the network vibrations in zeolites. The
210 first of the bands was assigned to Si-O stretching vibrations, the second to T-O
211 symmetrical stretching vibrations and the third to the vibrations generated by the T-O-T
212 bonds in aluminosilicates, probably with a four tetrahedral ring (Dutta and Barco,
213 1988). The gismondine observed in the XRD and SEM analyses were a likely source of
214 such vibrations. The spectrum also contained a wide signal at 360 cm^{-1} to 316 cm^{-1}
215 assignable to Ca-O vibrations. The wide signal at 860 cm^{-1} to 825 cm^{-1} appearing in
216 some types of zeolites is due to the Si-O vibrations in their terminal Si atoms. The
217 widening of the band peaking at 1087 cm^{-1} (calcite) might denote the presence of
218 amorphous calcium carbonate (not identified by XRD) or of carbonate phases such as
219 monocarboaluminate.

220

221 **3.4 ^{29}Si and ^{27}Al NMR-determined mineral phases**

222 The ^{27}Al MAS NMR spectrum for the 15 year MK/lime sample reproduced in Figure 7a
223 showed a very intense tetrahedral Al signal (54.9 ppm) with a pentahedral Al shoulder
224 (25 ppm to 45 ppm) and a second signal attributed to octahedral Al (6.4 ppm). Since

225 according to Richardson et al. (2010) Q^3 bridging sites across the interlayer generate a
226 resonance peaking at around 58 ppm, the signal at 54.9 ppm was attributed to
227 tetrahedrally coordinated aluminium in aluminosilicate (gismondine) (Khodabandeh and
228 davis, 2006). Anderson et al. (2006) identified a low-frequency peak at 35 ppm as an
229 aluminate phase associated with an amorphous or scanty crystalline aluminate
230 hydroxide, which would account for the shoulder observed here. The presence of the
231 octahedral Al signal at 6.4 ppm denoted Al replacement of Si in both the tetrahedral and
232 octahedral sheets in the zeolite's layered silicate structure.

233 The ^{29}Si NMR spectrum (Figure 7b) exhibited two signals, one very intense peaking at
234 90 ppm and the other less intense, peaking at 94 ppm, both in the interval characteristic
235 of Q^3 units. No signals were observed in the Q^1 , Q^2 unit interval, which would be
236 indicative of the presence of C-S-H gels forming during cement hydration or the
237 pozzolanic reaction between MK and lime. That confirmed the Raman and XRD
238 findings suggesting C-S-H gel disintegration and the formation of a gismondine-like
239 zeolite.

240

241 **3.5 FT-IR determined mineral phases**

242 The bands at 1008 cm^{-1} and 439 cm^{-1} observed on the FTIR spectrum for the 15 year
243 MK/lime paste (Figure 8) were due to the internal vibrations in TO_4 tetrahedra (T = Al,
244 Si). The former was associated with T-O bond asymmetrical stretching vibrations and
245 the latter to T-O bond internal bending vibrations (Andersen et al., 2006). Both T-O
246 bonds were present in the mineral phases identified. The intermediate intensity signal
247 peaking at 580 cm^{-1} was assigned to zeolite external ring vibrations.

248 The IR spectrum also exhibited a very intense signal peaking at 1470 cm^{-1} attributed to
249 CO_3 ν_3 asymmetric stretching vibrations generated by the carbonates in aragonite. The
250 signals at 874 cm^{-1} and 855 cm^{-1} were due to CO_3 ν_2 asymmetric bending vibrations

251 induced by calcite and aragonite, respectively. Similarly, the signals at 712 cm^{-1} and
252 700 cm^{-1} were respectively attributed to calcite and aragonite CO_3 ν_1 symmetric
253 bending (Gunasekaran et al., 2006). The presence of vaterite, which according to the
254 XRD findings accounted for 6 % of the sample, was detected as a shoulder at 745 cm^{-1} ,
255 confirming its low concentration. FT-IR likewise confirmed zeolite and carbonate
256 polymorph formation. The signals at 580 cm^{-1} and 438 cm^{-1} were associated with AlO_6
257 units, possibly indicative of a monocarboaluminate hydrate phase, which would
258 corroborate the Micro-Raman data.

259 **3.6 TG/DTA-determined mineral phases**

261
262 The TG thermogram for the paste (Figure 9) showed a total loss of 30.41 % across the
263 test interval ($25\text{ }^\circ\text{C}$ to $1000\text{ }^\circ\text{C}$), in which two substantial areas of weight loss could be
264 clearly identified. A loss of 9.34 % of the total was recorded in the lower ($25\text{ }^\circ\text{C}$ to
265 $300\text{ }^\circ\text{C}$) and of 16.84 % loss in the higher ($635\text{ }^\circ\text{C}$ to $789\text{ }^\circ\text{C}$) temperature range. The
266 DTA curve clearly showed five endothermic peaks in the former, the most intense of
267 which at $146.7\text{ }^\circ\text{C}$ and the four others at $50\text{ }^\circ\text{C}$, $93\text{ }^\circ\text{C}$, $110\text{ }^\circ\text{C}$ and $231\text{ }^\circ\text{C}$. The $50\text{ }^\circ\text{C}$
268 endothermic peak represented the loss of adsorbed moisture while the others were
269 associated with dehydroxylation of the mineral phases present in the 15 years sample.
270 According to the literature (Frías, 2006b; Foldvari, 2011), the main hydrated phases
271 resulting from the MK/lime pozzolanic reaction dehydroxylate in this temperature
272 interval: C-S-H gel at $110\text{ }^\circ\text{C}$ to $140\text{ }^\circ\text{C}$, C_2ASH_8 at $140\text{ }^\circ\text{C}$ to $160\text{ }^\circ\text{C}$, C_4AH_{13} at $170\text{ }^\circ\text{C}$
273 to $200\text{ }^\circ\text{C}$, carboaluminate at $200\text{ }^\circ\text{C}$ to $230\text{ }^\circ\text{C}$.

274 The presence of feldspar and zeolites, primarily gismondine, was confirmed by the
275 endothermic peaks at $88\text{ }^\circ\text{C}$ to $115\text{ }^\circ\text{C}$, $115\text{ }^\circ\text{C}$ to $208\text{ }^\circ\text{C}$ and $208\text{ }^\circ\text{C}$ to $290\text{ }^\circ\text{C}$,
276 representative of the consecutive loss of water molecules (0.5, 1.5 and 0.5 moles,

277 respectively). Given endothermal peak overlapping, they could not be clearly assigned
278 to any specific mineral phase.

279 The second interval exhibited a wide endothermal signal peaking at 762 °C, typical of
280 carbonate decarbonation in the CaCO₃ polymorphs identified by XRD, i.e., primarily
281 calcite, aragonite and some vaterite. The presence of carbonates was related to
282 carbonation of the portlandite in the MK/Ca(OH)₂ system over the 15 year curing,
283 storage and testing period.

284

285 **4. DISCUSSION**

286 Stratlingite, C₄AH₁₃, hydrogarnet and C-S-H gels are the majority short- and medium-
287 term hydrated phases in pozzolanic reactions taking place at 60 °C (Martínez and Frías,
288 2011; Frías and Cabrera, 2002; Frías and Sánchez de Rojas, 2003; Frías et al., 2013).
289 According to Silva et al. (1993)¹, hexagonal phases C₂ASH₈ and C₄AH₁₃ are
290 metastable, evolving under high temperature and humidity conditions toward a stable,
291 hydrogarnet family phase with variable compositions. This study showed for the first
292 time that exposure time plays a significant role even for stable mineral phases, a finding
293 not experimentally proven to date. Over the 15 years of exposure, all the (stable and
294 metastable) hydration phases evolved toward carbonates, feldspars and zeolites due to
295 degradation and heterogeneous nucleation.

296 Hydrogarnets, part of the nesosilicate family, constitute an isostructural group in the
297 cubic system with the general formula (SiO₄)₃A₃B₂, where A represents divalent cations
298 such as calcium, magnesium, iron or manganese and B trivalent cations such as
299 aluminium, iron or chromium. Hydrated hydrogarnets also contain water molecules in
300 their structure in the form of (OH)⁴⁻ groups, with the four hydrogen atoms replacing a
301 silicon atom in the (SiO₄)⁴⁻ tetrahedron. The crystalline divalent cations exhibit

302 dodecahedral, the trivalent cations octahedral and the silicon tetrahedral coordination.

303 Calcium silicate hydrate (C-S-H), strätlingite ($2\text{CaO}\cdot\text{Al}_2\text{O}_3\cdot\text{SiO}_2\cdot 8\text{H}_2\text{O}$) and a

304 hydrogarnet ($3\text{CaO}\cdot\text{Al}_2\text{O}_3\cdot 6\text{H}_2\text{O}$)-katoite ($3\text{CaO}\cdot\text{Al}_2\text{O}_3\cdot\text{SiO}_2\cdot 4\text{H}_2\text{O}$) solid solution may

305 co-exist in cements.

306 In strätlingite, aluminium atoms are tetrahedrally or octahedrally coordinated and

307 silicon atoms exist predominantly as Q^2 , Q^2 (1Al) or Q^2 (2Al) units. In 1976, Kuzel

308 (1976) described strätlingite structure based on the formula: $[\text{Ca}_2\text{Al}^{\text{VI}}(\text{OH})_6]$

309 $[\text{Al}^{\text{IV}}\text{SiO}_3(\text{OH})_2(\text{H}_2\text{O})_4]$, indicating that the $\text{Al}^{\text{IV}}/\text{Al}^{\text{VI}}$ ratio was equal to 1. Muller et al.

310 (1986) in turn determined the silicon in strätlingite to constitute a Q^3 (3Al) unit.

311 Strätlingite structure may be affected by the presence of alkalis, which significantly

312 reduce the $\text{Al}^{\text{IV}}/\text{Al}^{\text{VI}}$ ratio. This mineral consists in a brucite-like octahedral layer

313 comprising calcium and aluminium with an ideal composition of $[\text{Ca}_2\text{Al}(\text{OH})_6\cdot 2\text{H}_2\text{O}]^+$

314 and a double tetrahedral layer with ideal composition $[(\text{T},\square)_4(\text{OH},\text{O})_8\cdot 0.25\text{H}_2\text{O}]$.

315 Rinaldi et al. 1990) observed that the tetrahedral positions could be occupied by silicon

316 or aluminium ions and that 45 % of those positions in the double tetrahedral layer were

317 vacant. The calcium atoms were hepta-coordinated and occupied cationic positions

318 associated with the octahedral aluminium layer. The water molecules, most found in the

319 octahedral layer and projecting into the centre of the rings comprising the double

320 tetrahedral layers, hold the tetra and octahedral layers together by forming hydrogen

321 bonds. One additional water molecule is located in the centre of each tetrahedral ring.

322 The dimensions of the unit cell infer overlapping in three octahedral layers and three

323 interlayers (double tetrahedral layers) in direction [001] (Rinaldi, 1990). The degree of

324 structural disorder in the strätlingite microstructure forming in the cement pastes, which

325 depends essentially on the hydration conditions (Santacruz et al., 2016), must have

326 been substantial, for greater disorder would favour degradation.

327 The loosely organised structure of C-S-H gel hinders its characterisation with XRD. It
328 has nonetheless been shown to have a laminar crystalline structure similar to two
329 minerals: tobermorite, when the Ca/Si ratio ranges from 0.66 to 1.5, and jennite, when it
330 is over 1.5. The nanometer-scale laminar thickness (Santacruz et al., 2016) of this C-S-
331 H gel affects the mechanical and engineering properties of cements and concretes
332 Ioannidou et al., 2016). When the Ca/Si ratio = 0.6, C-S-H gel structure consists in a
333 double central sheet of six or seven calcium ions linked to oxygen ions connected on
334 each side to chains of silica tetrahedra. The silicate chain adopts a dreierketten
335 arrangement, i.e., a repeating unit consisting in three SiO₄ tetrahedra, one bridging and
336 two pairing. The latter two share oxygen atoms with a sheet of octahedral CaO while
337 the former does not. All the unattached oxygen atoms in the silicate chain are
338 protonated as hydroxyls. The interlayer space is occupied by water molecules and, if the
339 Ca/Si ratio is > 0.83, by calcium ions (Bonaccorsi et al., 2004). The main difference
340 between C-S-H gel and tobermorite / jennite is that in the former the chains are finite
341 and in the latter two infinite. The evolution of C-S-H gel structure is governed by the
342 Ca/Si ratio. If the ratio is 1.5, the protons in the bridging SiO₄ tetrahedra in the
343 dreierketten units are progressively replaced with calcium ions, which are taken up in the
344 hydrated interlayer space. If the Ca/Si ratio declines to values of around 1.0, the degree
345 of chain polymerisation declines and the hydroxyl groups are ionised, whereby the
346 negative charge in the laminae is maintained and the calcium ions in the interlayer
347 remain unaffected (Nachbaur et al., 1998).

348 The structures of the materials described may be deemed to consist in mixes of surface
349 groups. The calcium ion concentration and pH in the solution are known to ultimately
350 induce calcium ion supersaturation (Madani et al., 1990). Four types of surface groups
351 may form: SiOH, SiO⁻, SiOCa⁺ and SiOCaOH, the first two at low and the last two at

352 higher Ca concentrations. The latter condition induces high proportions of reactive
353 SiOCa^+ positions. Such differential surface reactivity depending on pH and especially
354 Ca availability might explain the present observations: the ready nucleation and growth
355 of calcium feldspars and gel fibres on low-calcium substrates (with SiOH and SiO^-
356 surface groups) and the formation of calcium carbonates and zeolites on high-calcium
357 substrates (SiOCa^+ and SiOCaOH).

358 Zeolites may form via hydrothermal synthesis of MK and lime in the presence of high
359 concentrations of tetra- and penta-coordinated aluminium and a minimum content of
360 hexa-coordinated Al (Ferro et al., 2002). The presence of three forms of coordinated
361 aluminium observed with NMR might explain zeolite nucleation and growth on a
362 substrate consisting in high-calcium conglomerate. Such conditions would suffice for
363 the formation of zeolitic calcium (gismondine), a double chain zeolite with four
364 interconnected rings exhibiting slightly deformed P21/c symmetry and tetrahedra with
365 centres alternatively occupied by Si or Al. Ca in turn is located on one of the four
366 bending positions that connect two tetrahedral oxygens to water molecules, distributed
367 across six positions (Fischer and Schramm, 1971).

368 The sketch in Figure 7 shows how gismondine would nucleate via the cycling of double
369 tetrahedral chains resulting from the disintegration of the initial mineral phases
370 (strätlingite, C_4AH_{13} , C_2ASH_8 , hydrogarnets and C-S-H gel). In Si gismondine the Al
371 distribution is orderly and only one type of extra-framework cation (calcium) is present
372 in the channels. Extra-framework cations are a determinant in the mechanism governing
373 zeolite deformation (Madani et al., 1990), while polyhedral coordinated Ca serves as a
374 template that also drives framework deformation (Fois et al., 2005; Betti et al., 2007).

375 Zeolite formation as observed in cementitious systems has also been described in
376 cement- and lime-free media, such as in geopolymer synthesis, where it occurs along

377 with amorphous phase generation (Davidovits, 2008). The chemical composition of
378 geopolymers is similar to that of C-S-H gel and zeolite-like materials. Some authors
379 have described geopolymers as zeolite precursors (Liew, 2016). Zeolite crystallization is
380 favoured by high water content, high curing temperature and lengthy curing periods
381 (Provis, 2005). The number of zeolite crystallites also grows with time (Duxson et al.,
382 2005).

383 The results of this study show that exposing lime- and pozzolan-based cementitious
384 matrices to high temperatures for long curing times induces mineralogical
385 transformations that primarily induce zeolite formation. That process, due to a high
386 calcium concentration and high pH that favour the formation of SiOCa^+ - and SiOCaOH -
387 like, zeolite-seeding surface groups, entails the cycling of the double chain in C-S-H gel
388 and concomitant gismondine formation.

389 The generally accepted notion that metastable phases may co-exist at high curing
390 temperatures and evolve thermodynamically toward a stable phase may therefore be
391 true, albeit only partially, for the experimental trials conducted to date cover relatively
392 short hydration periods. The present results reveal that over much longer curing times
393 such phases practically disappear and give way to others, suggesting new lines of
394 research for the international scientific community.

395 **5. CONCLUSIONS**

396 In this work the reaction between metakaolinite and calcium hydroxide at 60°C after 15
397 years was studied. Although it is known that stable and metastable phases coexist at
398 times less than 5 years, the formation of gismondine-type zeolites occurs at longer
399 hydration times.

400 Microstructurally, aggregates from the degradation of the initial compounds C-S-H,
401 C_4AH_{13} , stratlingite and hydrogarnets were observed, which serve as crystallization

402 substrates for new phases such as calcite and mainly calcium zeolites of the gismondine
403 type. The presence of these phases is confirmed by characterization techniques such as
404 XRD, Micro-Raman, FT-IR and ^{27}Al and ^{29}Si NMR.

405 The formation of calcium zeolite, gismondine, is postulated that occurs by cycling
406 double chains of the initially formed C-S-H gel. This shows the formation of zeolites in
407 systems with lime, as currently has only been described in geopolymeric systems, with
408 high alkali content.

409 This new understanding will have a direct impact on the manufacture of metakaolinite-
410 blended cements and their engineering properties, for the structures engineered today
411 are designed for a future service life of 100 to 200 years.

412

413 **ACKNOWLEDGEMENTS**

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Figure captions

Figure 1. Device used for 15 year curing

Figure 2. XRD spectrum for 15 year paste

Figure 3. SEM images: a) aluminate crystals and b) strätlingite crystals

Figure 4. SEM images: a) fibres; b) and c) feldspar forming on conglomerate; d) calcite crystals

Figure 5. Calcium zeolites forming on conglomerate (a and b)

Figure 6. Micro-Raman spectrum of the sample (laser $\lambda = 785$ nm); C = calcite; Z = zeolite.

Figure 7. NMR spectrum for the 15 year MK/lime sample: a) ^{27}Al MAS and b) ^{29}Si MAS NMR.

Figure 8. FT-IR spectrum for the 15 year paste

Figure 9. TG/DTA thermogram for the 15 year paste

Figure 10. Gismondine formation by cycling of C-H-S double chains

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629

Table[Click here to download Table: Tables.docx](#)**Table 1. Chemical composition of raw MK (%)**

Oxide	SiO ₂	Al ₂ O ₃	Fe ₂ O ₃	CaO	MgO	SO ₃	Na ₂ O	K ₂ O	TiO ₂	LOI
MK	51.60	41.60	4.64	0.09	0.16	-	0.01	0.62	0.83	0.60

Oxide (%)	C ₄ AH ₁₃	C ₂ ASH ₈	Conglomerates	Fibre	Gismondine	Garronite	Feldspar
Na ₂ O	n.d.	1.43±0.39	3.49±0.55	5.28±0.82	1.32±0.56	1.87±0.69	1.53±0.79
MgO	n.d.	n.d.	0.64±0.08	n.d.	n.d.	n.d.	n.d.
Al ₂ O ₃	22.92±0.82	25.12±2.27	29.34±1.79	11.72±0.94	38.00±1.96	33.10±2.51	33.29±2.35
SiO ₂	1.43±0.12	13.66±1.58	51.24±0.87	17.44±0.65	40.94±2.18	47.3±2.36	47.75±1.28
K ₂ O	n.d.	1.46±0.72	1.99±0.38	n.d.	n.d.	n.d.	0.75±0.12
CaO	75.65±1.27	58.33±1.27	13.30±0.72	65.56±0.41	19.74±1.64	17.66±1.29	16.68±0.69
Si/(Si+Al)	-	-	-	-	0.51	0.58	-

Table 2. SEM/EDX analysis of the 15 years paste

n.d. = not detected

Figure 1
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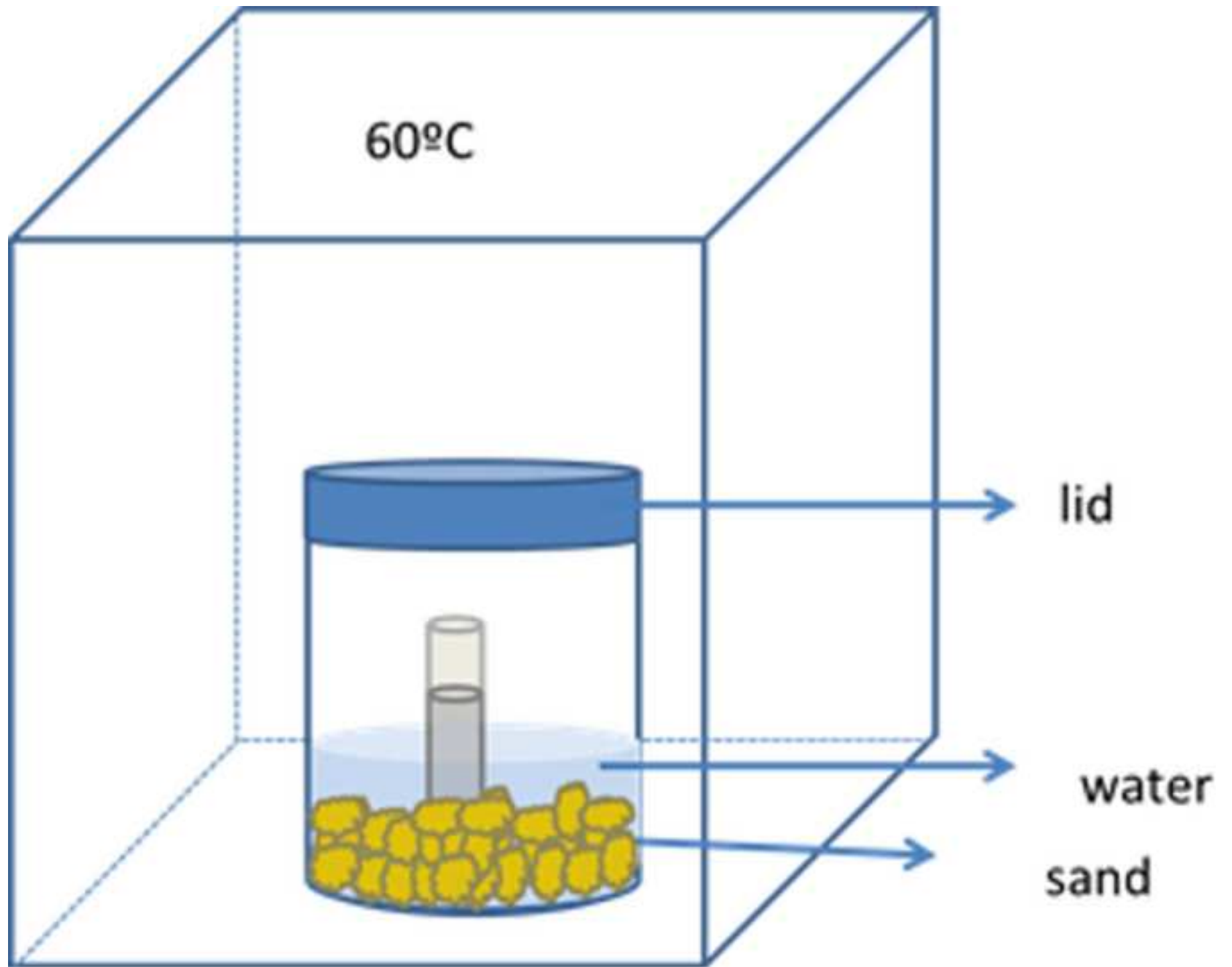


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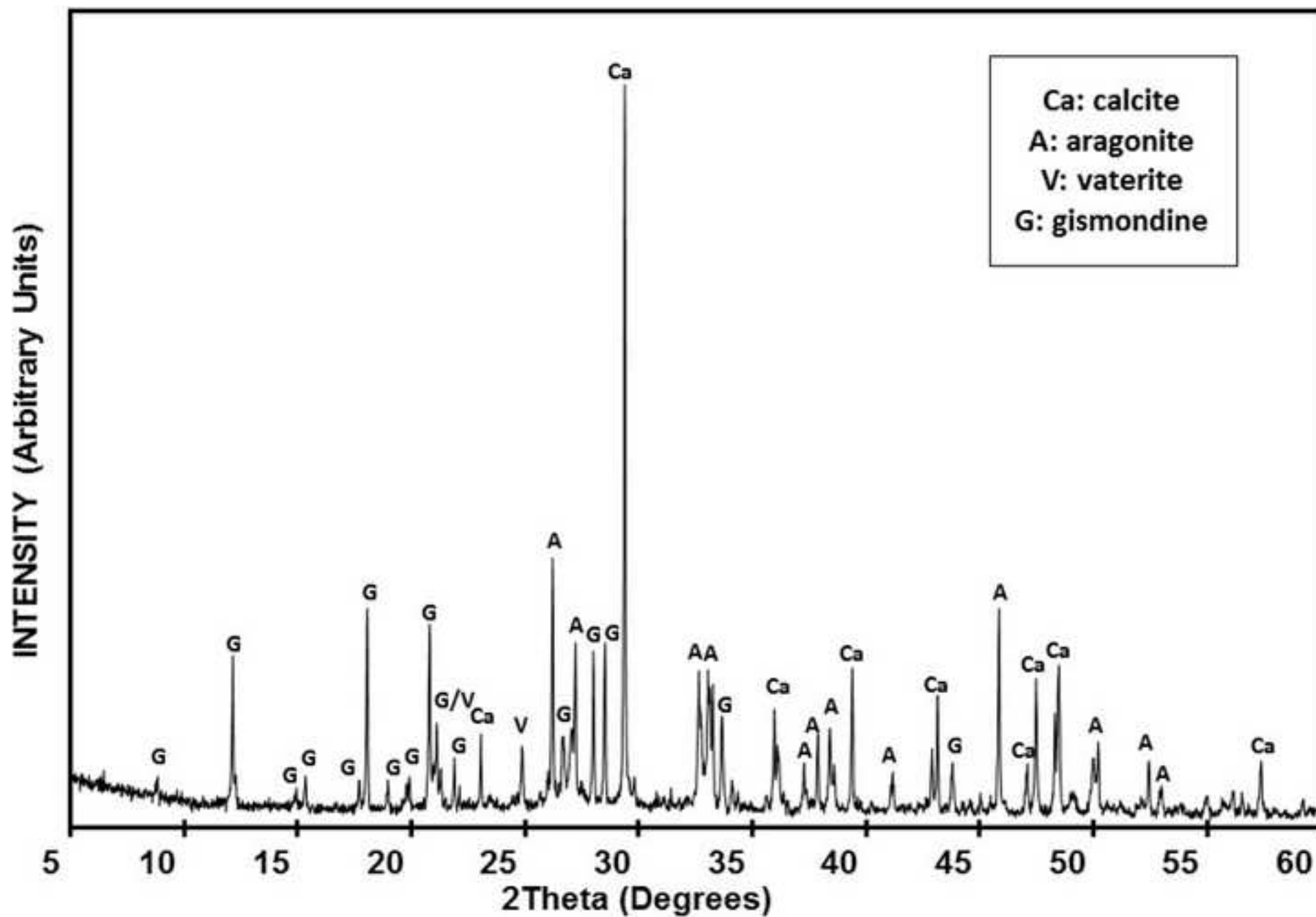


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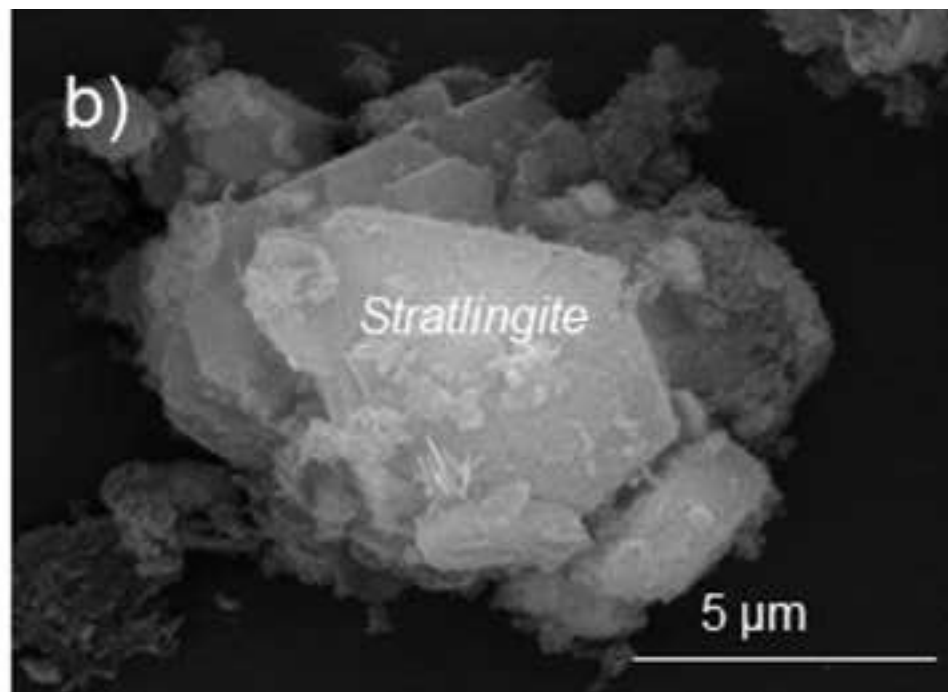
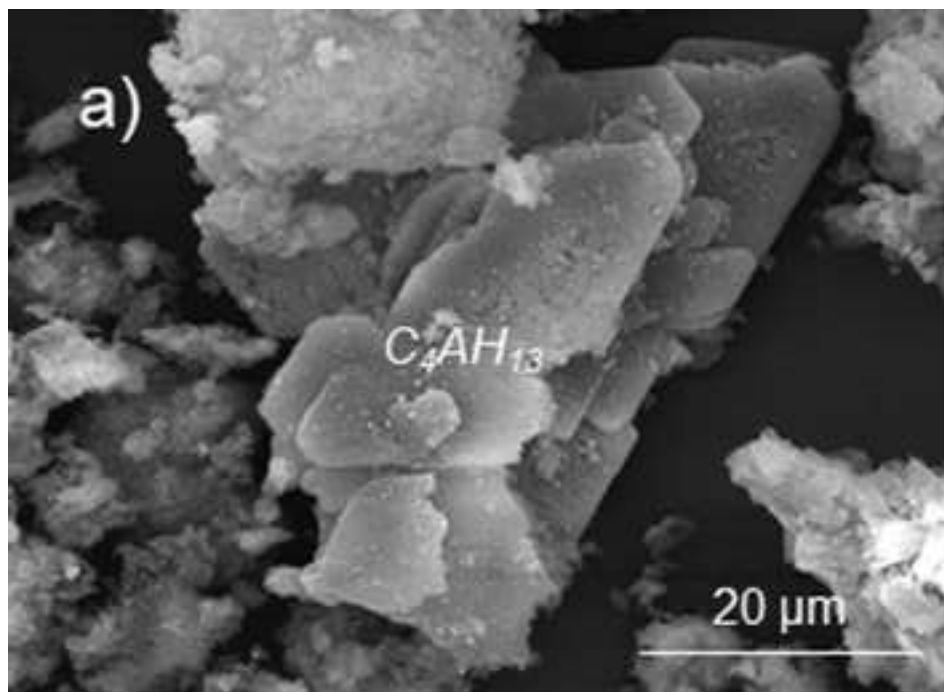


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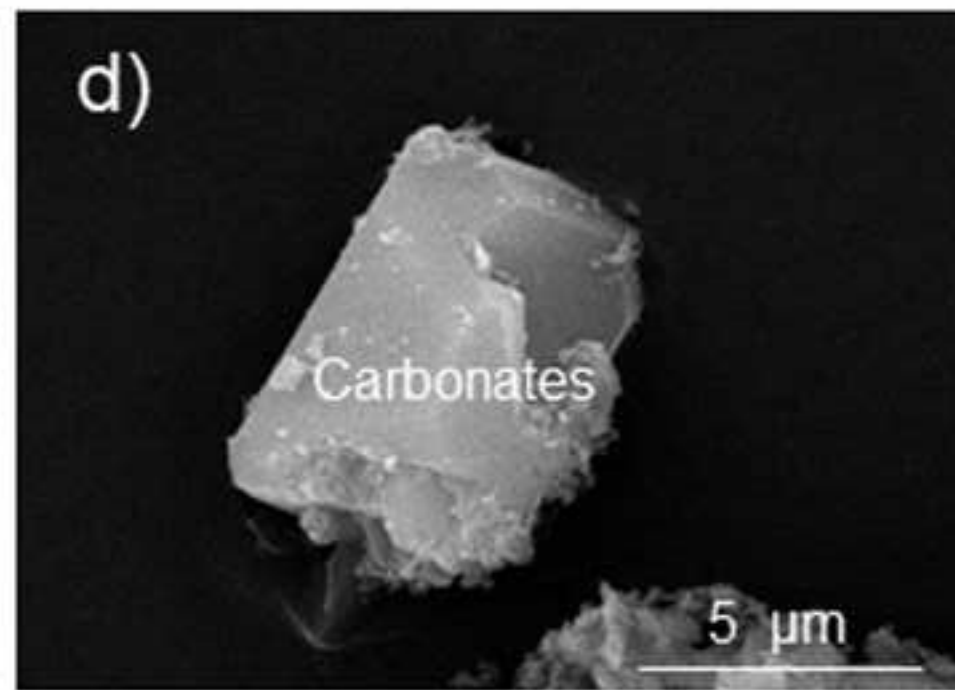
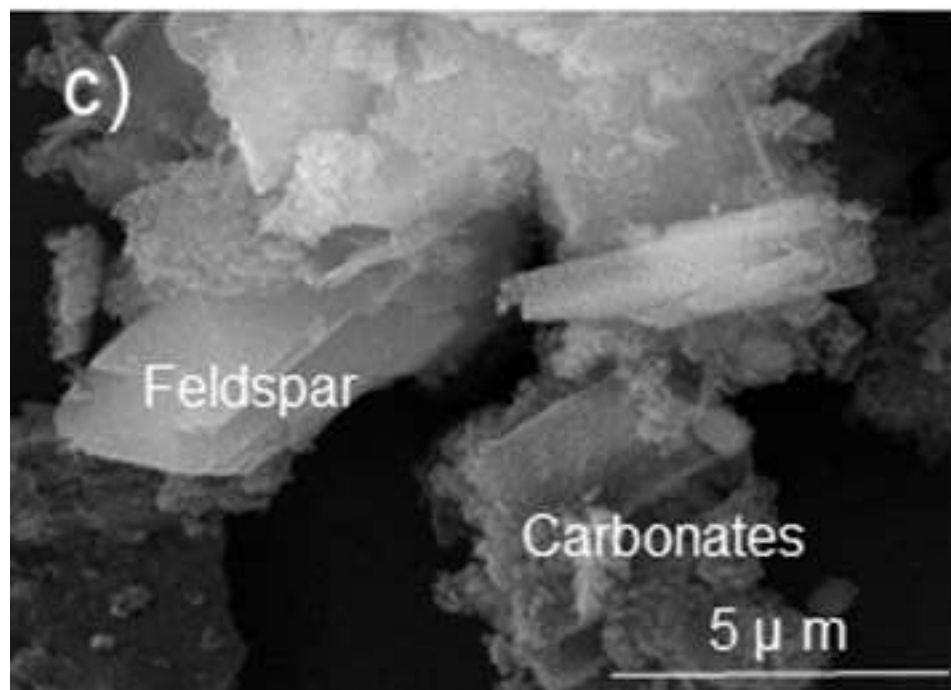
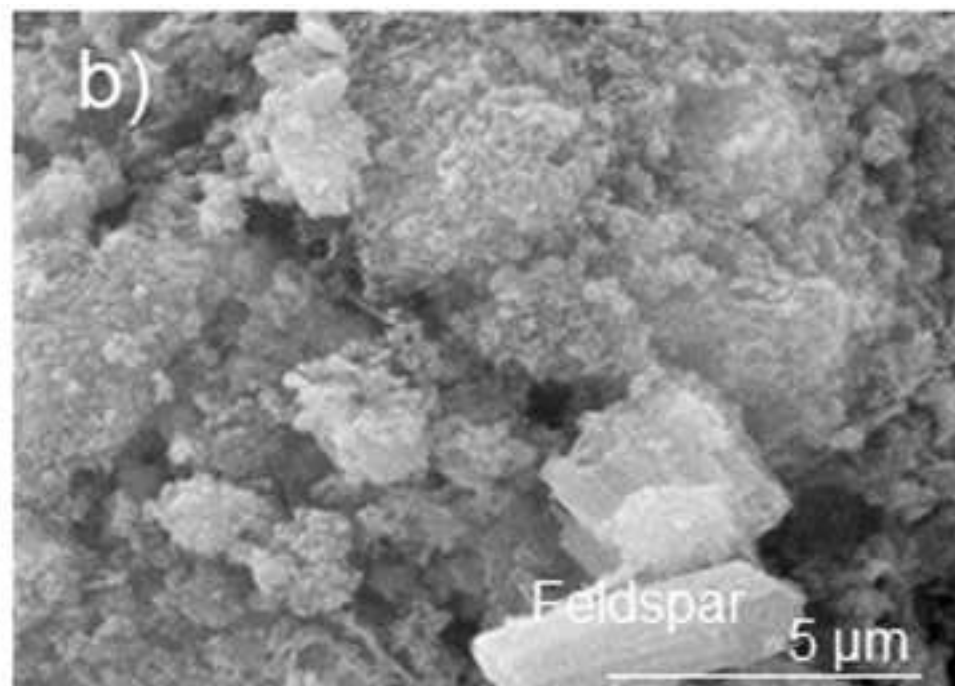
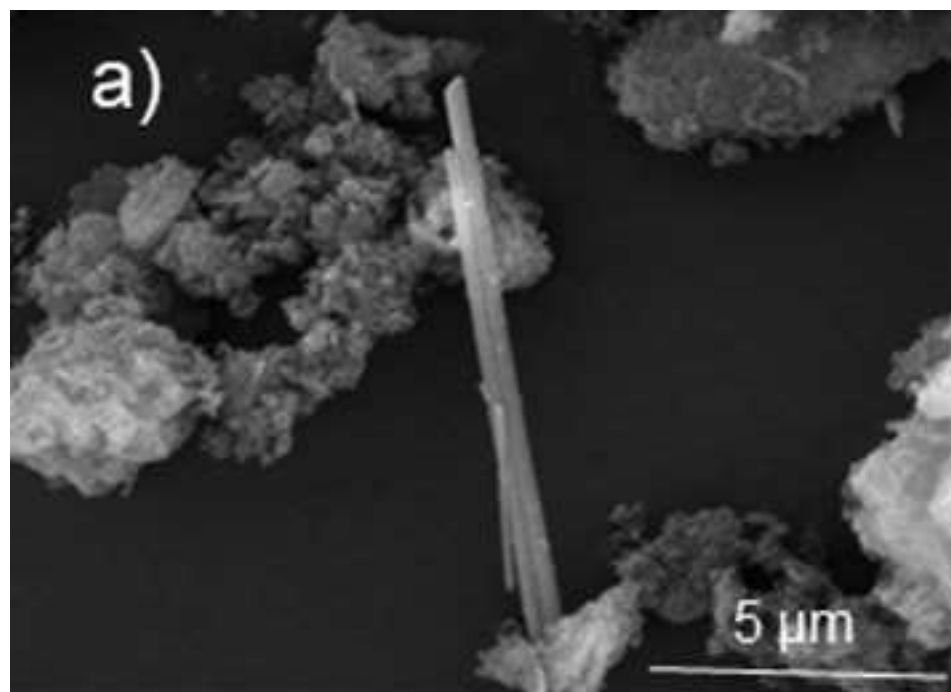


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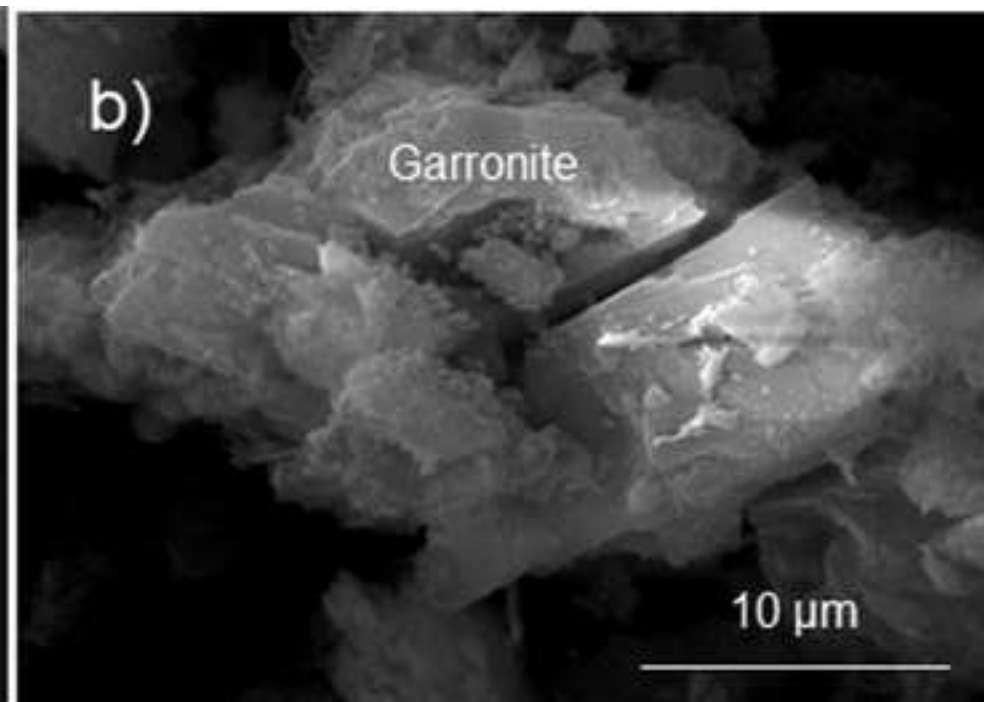
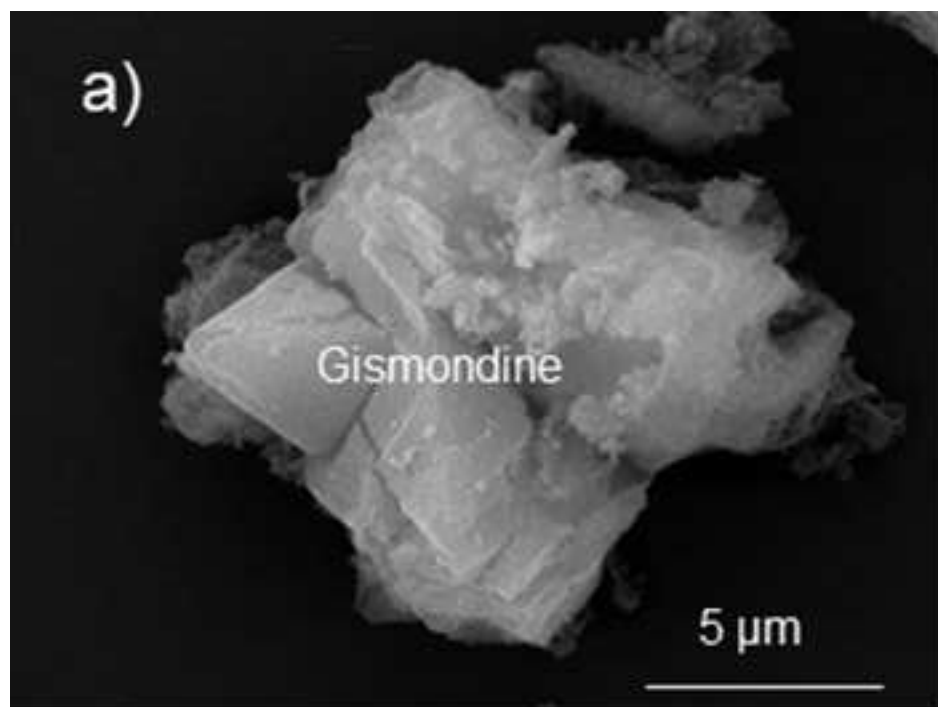


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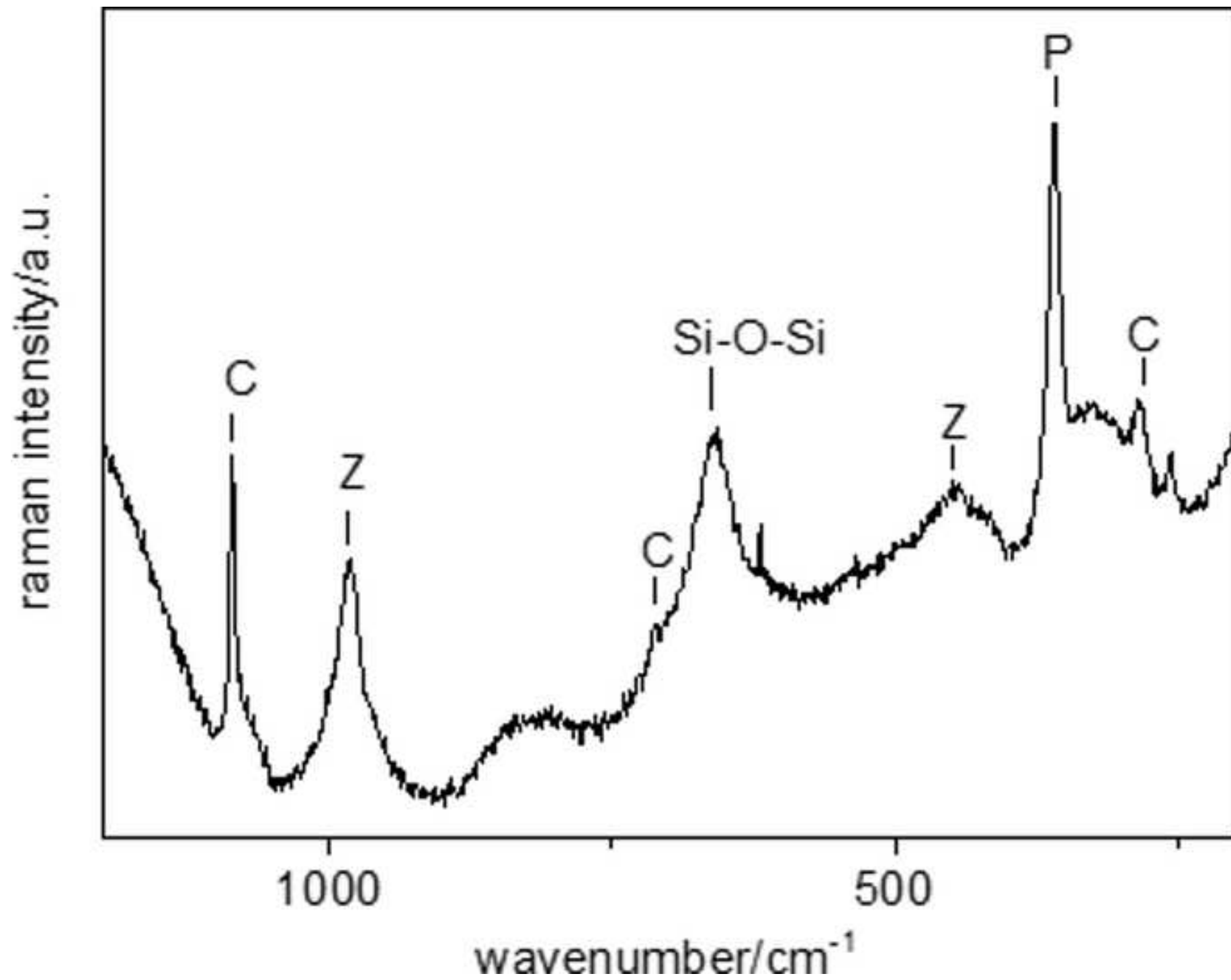


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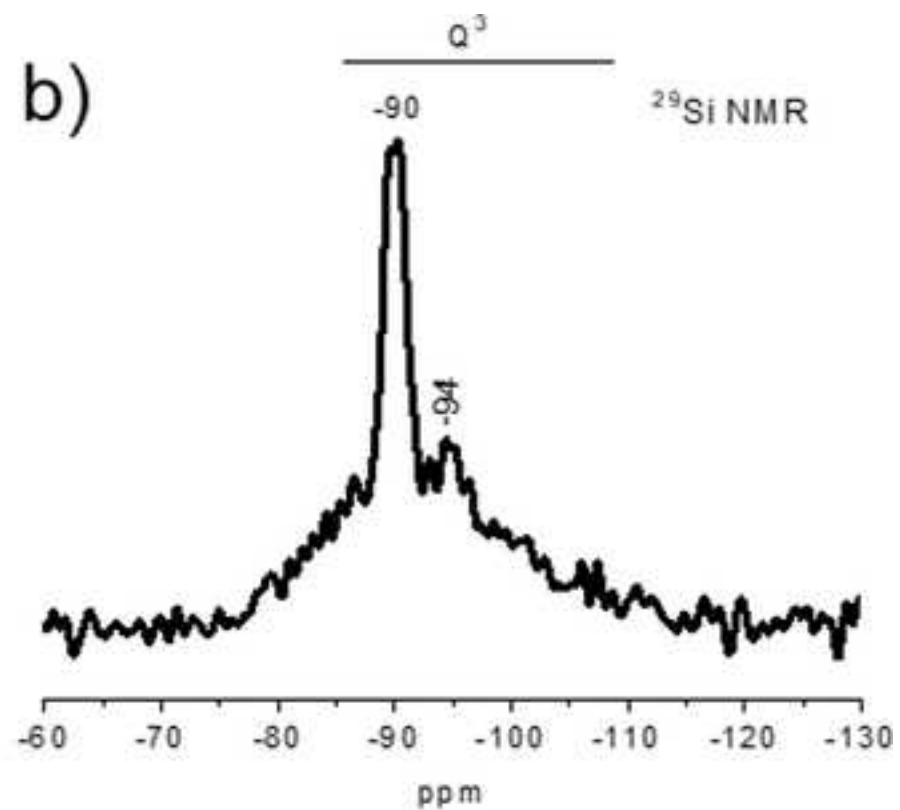
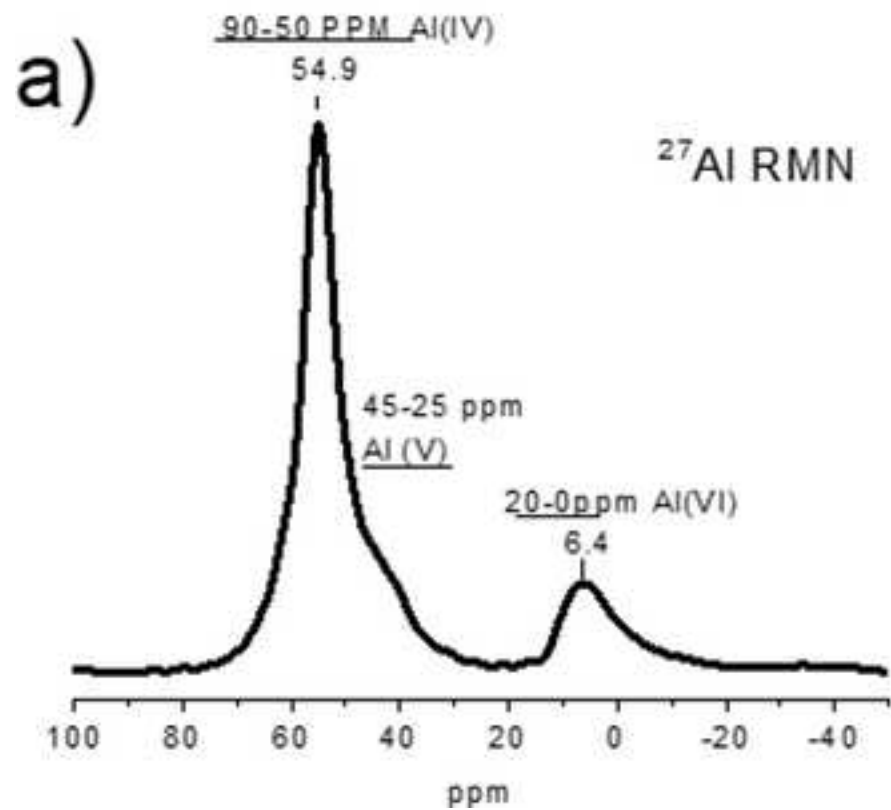


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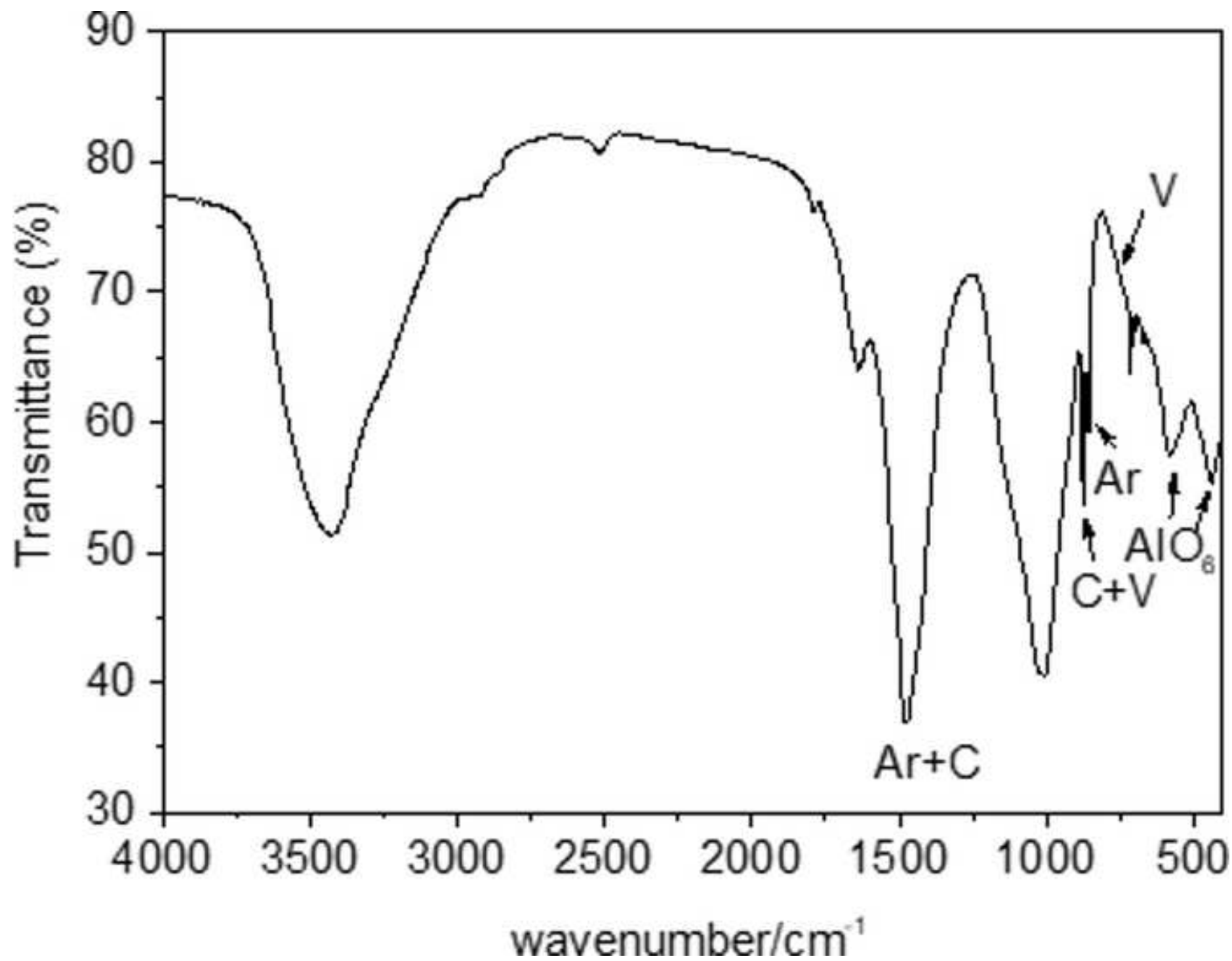


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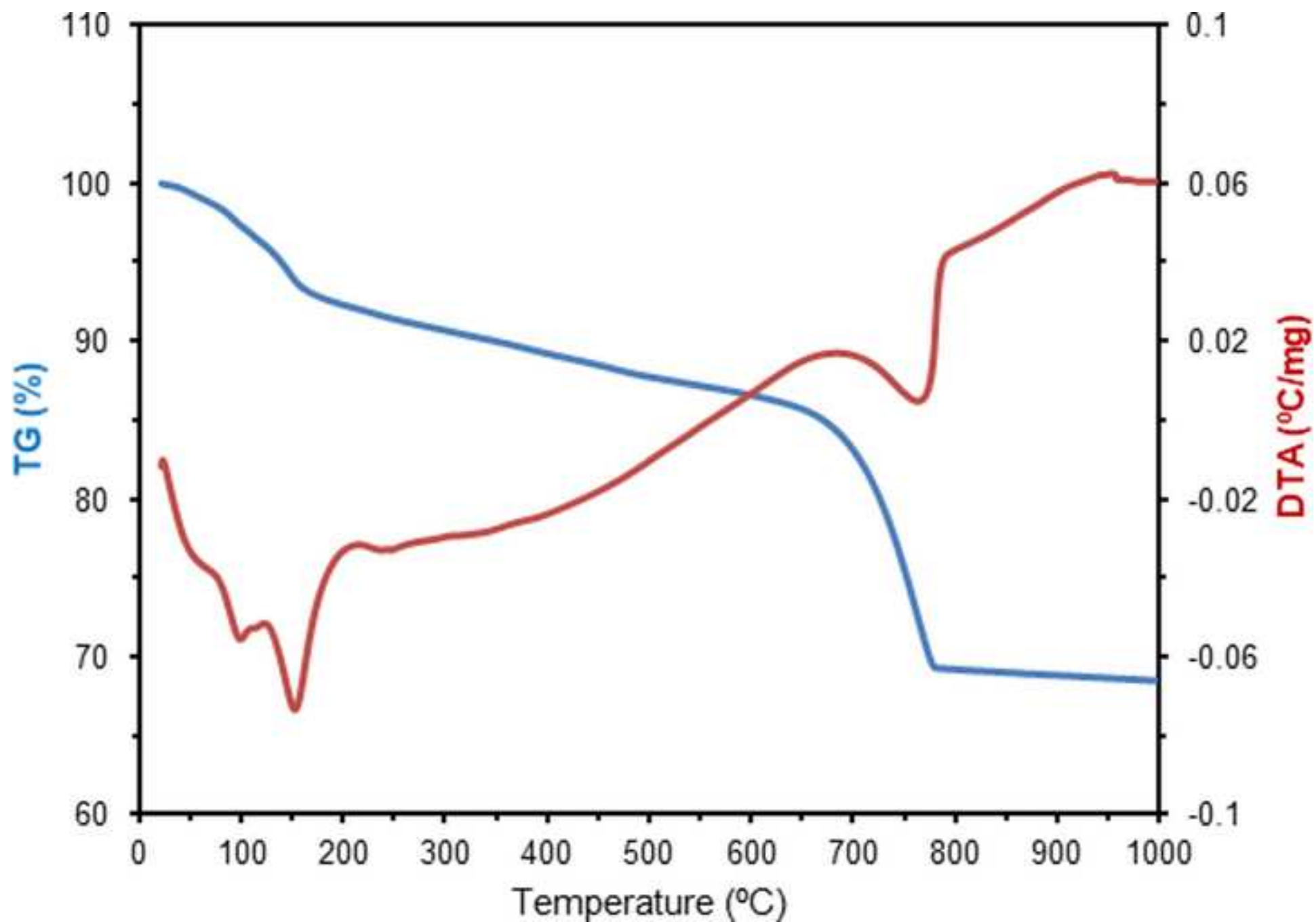


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